DROPS STONES RESULTING FROM SNOW-AVALANCHE DEPOSITION ON LAKE ICE

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ABSTRACT. Dirty snow avalanches have been observed to carry considerable amounts of rock debris on to lake ice at the foot of scree slopes. As ice breaks up in the spring thaw, this material is carried back and forth on ice floes and is gradually deposited in the lake. In some areas this produces typical drop stones of rock debris in predominantly fine-grained deposits. Most avalanche debris is very angular which enables avalanche drop stones to be differentiated from those of glacial or other drift-ice origins. However, where avalanches incorporate glacial debris, such deposits may be indistinguishable from those formed by floating glacier ice.

Résumé. Souppoudrage de sédiments résultant des apports d’avalanches de neige sur la glace de lac. On observe que les avalanches de neige sale transportent une quantité considérable de débris rocheux jusque sur la glace de lac au pied des pentes d’éboulis. Lorsque la glace se brise au moment de la fusion printanière, le matériel est transporté et est déposé graduellement dans le lac. Dans certaines zones ce produit un souppoudrage de sédiment caractéristique fait de débris rocheux dans des dépôts où dominent les matériaux fins. Beaucoup d’apports d’avalanches sont très anguleux ce qui peut permettre de différencier les sédiments souppoudrés issus d’avalanche de ceux dont l’origine est un glacier ou une congère construite par le vent. Cependant, lorsque les avalanches ont incorporé des débris glaciaires, leurs dépôts sont impossibles à distinguer de ceux formé par des fragments flottants de glace de glacier.


The geomorphic significance of snow avalanches in Arctic and alpine areas has been underestimated. Under favourable conditions, especially when travelling over surfaces of loose, unconsolidated rock fragments, avalanches may incorporate large amounts of rock debris (Rapp, 1960; Gardner, 1970; Luckman, 1972). Continued avalanche activity in such locations leads to the production of well-developed landforms known as avalanche boulder tongues (Rapp, 1959; Luckman, unpublished). Usually the avalanche debris is deposited within the terminal zone of the track but where avalanches extend on to ice-covered water bodies there may be considerable redistribution of these deposits.

Annual debris accumulation from rock falls and snow avalanches has been recorded on seven scree slopes in Surprise Valley, Jasper National Park, since 1968 (Luckman, 1971, unpublished). Two of the sites studied flank lakes so that in several cases avalanche deposition has extended on to the lake ice. When this ice breaks up, the rock debris, ablated from the avalanche snow, is carried by the ice floes over the lake basin and subsequently deposited as the ice melts (Fig. 1). This secondary deposition may be concentrated in one location or dispersed over the lakes depending on the manner of debris release (tilting, overturning or gradual ablation) and the pattern of drift-ice movement in response to the winds or lake currents.

The avalanche deposits themselves may range from pure snow to a dirty snow and rock admixture. The characteristics of the rock debris depend on its origin. The most common debris sources are scree slopes or loose debris swept from the cliff zone of the avalanche track. These materials are a characteristically heterogeneous, poorly sorted mixture of angular, often freshly broken fragments ranging from boulder to silt size (see Fig. 1). Since the rock debris can be derived from any unconsolidated material in the avalanche track, it could also include fluvial or glacially moulded debris.

The major implications of these observations are three-fold.

(a) Although these specific observations only relate to two small alpine lakes over a relatively short time period (6 years), morphological evidence suggests these processes have been active over a much longer period (several thousand years?), and similar deposits have been observed on larger water bodies
elsewhere in the Canadian Rockies, in Scandinavia (Rapp, 1960, fig. 48) and on sea ice in the Canadian Arctic (Bones, unpublished, pl. 14). This suggests that, over long periods of time, snow avalanches and other mass movements (rock falls, slush avalanches, etc.) could locally deposit significant amounts of debris on lake or near-shore ice. Estimates from debris-accumulation measurements on scree with considerable avalanche erosion indicate mean annual deposition of the order of 0.5–5 mm m$^{-2}$ year$^{-1}$ (see Luckman, unpublished, p. 273).

![Figure 1](image_url)

**Fig. 1.** Partially submerged ice floe on Lake Helen, Surprise Valley, Jasper National Park, 8 June 1970. The thick debris cover is typical of avalanche-incorporated scree material. The large boulder is about 0.20 m$^3$ (1.2 m long).

(b) The dispersal and melting of this debris-covered ice may result in the formation of “drop stones” in areas of otherwise predominantly fine-grained sediments (lacustrine or marine). Such deposits are commonly ascribed to the action of floating glacier ice or drift ice which derive their rock debris from glaciers, fluvial deposition on near-shore ice (Dangeard and Vanney [1974]) or ice-foot erosion of beach material (Dionne, 1974). Since the vast majority of avalanche drop stones are markedly angular, it should be possible to differentiate them by their morphology (absence of rounded, polished or striated debris and their locale (evidence of past or present avalanche activity near the shore). However, where avalanches have incorporated till or fluvial material this distinction cannot be made. An important corollary of this is that in some cases the presence of striated or polished drop stones is not conclusive evidence for floating glacier ice.

(c) The dispersal of avalanche deposition by lake ice may inhibit the development of typical avalanche landforms in the lowest part of avalanche tracks which terminate in lakes.

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REFERENCES


